

1 Diagnosis of multilayer clouds using photon path 2 length distributions

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4 Received 28 December 2009; revised 5 June 2010; accepted 11 June 2010; published XX Month 2010.

5 [1] Photon path length distribution is sensitive to 3-D cloud structures. A detection
6 method for multilayer clouds has been developed, by utilizing the information of photon
7 path length distribution. The photon path length method estimates photon path length
8 information from the low level, single-layer cloud structure that can be accurately
9 observed by a millimeter-wave cloud radar (MMCR) combined with a micropulse lidar
10 (MPL). As multiple scattering within the cloud layers and between layers would
11 substantially enhance the photon path length, the multilayer clouds can be diagnosed by
12 evaluating the estimated photon path information against observed photon path length
13 information from a co-located rotating shadowband spectrometer (RSS). The
14 measurements of MMCR-MPL and RSS at the Atmospheric Radiation Measurement
15 (ARM) Southern Great Plains (SGP) site have been processed for the year 2000. Cases
16 studies illustrate the consistency between MMCR-MPL detection and the photon path
17 length method under most conditions. However, the photon path length method detected
18 some multilayer clouds that were classified by the MMCR-MPL as single-layer clouds.
19 From 1 year statistics at the ARM SGP site, about 27.7% of single-layer clouds detected
20 by the MMCR-MPL with solar zenith angle less than 70° and optical depth greater than 10
21 could be multilayer clouds. It suggests that a substantial portion of single-layer clouds
22 detected by the MMCR-MPL could also be influenced by some “missed” clouds or by the
23 3-D effects of clouds.

24 **Citation:** Li, S., and Q. Min (2010), Diagnosis of multilayer clouds using photon path length distributions, *J. Geophys. Res.*,
25 *115*, XXXXXX, doi:10.1029/2009JD013774.

26 1. Introduction

27 [2] Detailed knowledge of the radiative properties of
28 atmospheric constituents is crucial to properly characterize
29 climate forcing mechanisms and quantify the response of the
30 climate system. An important challenge is detecting the
31 three-dimensional (3-D) structure of clouds and aerosols,
32 and properly modeling the effects of this structure on radi-
33 ative transfer. This is essential to reduce ambiguity in the
34 retrieval of atmospheric properties and to improve radiative
35 parameterization in models. Current ability to resolve 3-D
36 cloud structure is limited to scanning pulsed active sensors
37 and imaging instruments. However, no single ground-based
38 sensor has proven to be capable of doing the job for all of
39 the wide variety of atmospheric cloud situations. In general,
40 the laser devices are excellent for detecting essentially all
41 clouds that are visible from the ground and are within the
42 instruments’ height range. The laser systems are unable to
43 provide any information about higher cloud layers when
44 lower liquid-water layers are present. The great strength of

radar is its ability to penetrate clouds and reveal multiple 45
layers aloft. Although its sensitivity is impressive, the 46
millimeter-wave cloud radar fails to detect some of these 47
clouds, especially if the clouds are composed of small 48
hydrometeors, or the clouds may be thinner than the radar 49
sample volume depth resulting in partial beam filling and 50
reduced reflectivity [*Clothetaux et al.*, 2000]. 51

[3] Information of “missed” cloud layer is extremely 52
important for the Broadband Heating Rate Profile (BBHRP), 53
since “missed” upper layer clouds would substantially 54
impact radiation heating profiles. Figure 1 shows the cal- 55
culated SW, longwave (LW), and total heating rates for a 56
single-layer cloud, a double-layer water cloud, and an ice 57
cloud over water cloud at solar zenith angle of 45°. For the 58
LW calculation, we used the U.S. standard atmospheric 59
profile. In the calculation of double-layer cloud cases, we 60
added a “missed” water or ice cloud layer with water path of 61
10 g/m² (cloud optical depth about 1) above the lower water 62
cloud layer and reduced the lower layer water cloud path to 63
190 g/m² to ensure the same total water path of 200 g/m² for 64
all cases. The SW reaching the surface for three cases are 65
124.1, 122.8, and 122.5 w/m², respectively, whereas the 66
upwelling SW at the TOA are 376.1, 377.5, and 379.5 w/m², 67
respectively. Clearly, the differences of SW at both bound- 68
aries with/without “missed” cloud layer are very small, 69

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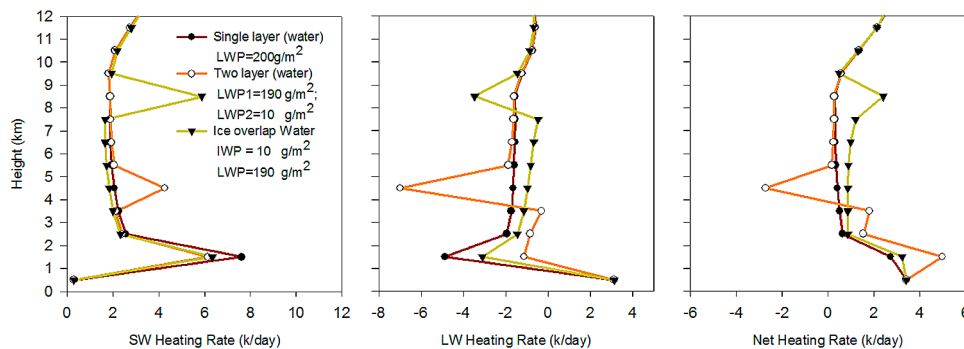


Figure 1. Broadband heating rate profile.

70 within the measurement uncertainty. However, the heating
 71 rate profiles are substantially different. Although a “missed”
 72 cloud layer does not occur all the time, statistical informa-
 73 tion of “missed” cloud layer is extremely valuable for
 74 BBHRP. Furthermore, this simple calculation reinforces
 75 that the radiation closure at the boundaries cannot ensure
 76 the accuracy of the heating profile. There is an urgent need
 77 to exploit other means to detect the 3-D structure of clouds
 78 and aerosols.

79 [4] For a long time, the remote sensing community has
 80 recognized the advantages of using the oxygen A band and
 81 has sought ways to exploit these advantages to measure
 82 atmospheric properties and constituents. Because oxygen is
 83 a well-mixed gas in the atmosphere, the pressure depen-
 84 dence (as a surrogate of altitude) of oxygen A band
 85 absorption line parameters provides a vehicle for retrieving
 86 photon path length distributions from spectrometry of the
 87 oxygen A band. The concept underlying oxygen A band
 88 retrievals is the principle of equivalence, which allows
 89 assessment of atmospheric radiative properties at any nearby
 90 wavelength from a photon path length distribution mea-
 91 surement at one particular band [Irvine, 1964; 1966; van de
 92 Hulst, 1980]. This is possible because the scattering prop-
 93 erties of cloud and aerosol vary slowly and predictably with
 94 wavelength and 760 nm is a useful central wavelength,
 95 reasonably representative of the entire solar shortwave.
 96 Photon path length distributions, a hidden property of
 97 standard radiation transfer models, are controlled by spatial
 98 distributions of scattering and absorption.

99 [5] Many efforts have been made to utilize photon path
 100 length distribution in oxygen A band as a tool in remote
 101 sensing [Grechko et al., 1973; Fischer and Grassl, 1991;
 102 Fischer et al., 1991; O’Brian and Mitchell, 1992; Harrison
 103 and Min, 1997; Pfeilsticker et al., 1998; Veitel et al., 1998;
 104 Min and Harrison, 1999; Portmann et al., 2001; Min et al.,
 105 2001; Min and Clothiaux, 2003; and Min et al., 2004; Min
 106 and Harrison, 2004; and many others]. In particular, Min
 107 and Clothiaux [2003] demonstrated that two independent
 108 pieces of information (mean and variance) are retrievable
 109 from a modest resolution Rotating Shadowband Spectrom-
 110 eter (RSS). Analysis of the variance and mean of the photon
 111 path length distribution from RSS measurements at the
 112 Atmospheric Radiation Measurement (ARM) Southern
 113 Great Plains (SGP) site illustrates how sensitive the photon
 114 path length distribution is to the cloud vertical profile. In this
 115 study, we further exploit the unique potential of photon path

length distribution to detect the 3-D structure of clouds and
 116 investigate how many clouds may be “missed” by the
 117 combination of a millimeter-wave cloud radar (MMCR) and
 118 a micropulse lidar (MPL) in a 1 year routine observation.
 119 Simply flagging possible “missed” clouds in routine
 120 MMCR-MPL observation is extremely valuable, as most
 121 ARM cloud products primarily use cloud retrievals from
 122 the MMCR.
 123

2. Methodology 124

2.1. Retrieval of Oxygen A Band Photon Path Length Distribution 125

[6] On the basis of the equivalent theory, the relationship
 127 between radiance measured in a spectral region free of the
 128 molecular absorption (such as at wavelengths outside the
 129 oxygen A band) to radiances measured within an absorption
 130 line can be written as 131

$$I_\nu = I_0 \int_0^\infty p(l) e^{-\kappa_\nu l} dl, \quad (1)$$

where I_0 and I_ν are radiances outside and within an
 132 absorption line, respectively, and $p(l)$ is the photon path
 133 length distribution. The transmission function $e^{-\kappa_\nu l}$ depends
 134 on the optical path length l and gaseous absorption κ_ν . The
 135 well-known effect of pressure broadening on line shape,
 136 which is a consequence of the dependency of κ_ν on pres-
 137 sure P and temperature T reveals information about the
 138 distribution of photon path length with pressure. The photon
 139 path length distribution can be derived from an inverse
 140 Laplace transform. Min and Clothiaux [2003] have devel-
 141 oped an approach to infer photon path length distributions
 142 from RSS measurements. This retrieval algorithm obtains
 143 empirical calibration coefficients of slit functions from
 144 clear-sky direct beam observations and applies them to
 145 diffuse irradiance measurements under cloudy sky condi-
 146 tions. Assuming $p(l)$ to be a simple γ distribution and using
 147 the existence of the Laplace transform, the photon path
 148 length distribution is retrieved from diffuse irradiance
 149 measurements. The detailed retrieval algorithm was pro-
 150 vided by Min and Clothiaux [2003]. More important, on the
 151 basis of the information content analysis and RSS perfor-
 152 mance, Min and Clothiaux [2003] also provided the
 153 assessment of uncertainty in both mean and variance esti-
 154

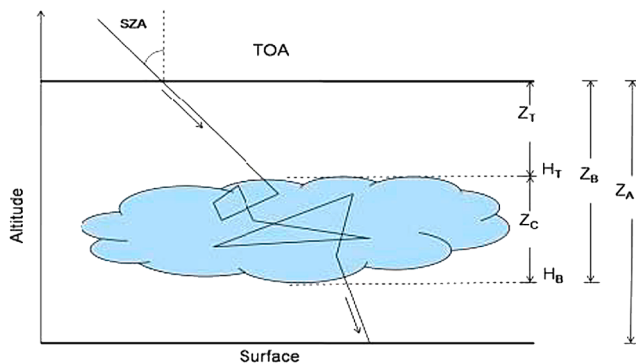


Figure 2. Schematic of photon path length in the atmosphere. H_T and H_B are the cloud top and base heights, respectively. Z_A , Z_T , and Z_B are the pressure-weighted oxygen cumulated paths for entire atmosphere, from TOA to the cloud top, and from TOA to the cloud base, respectively. Z_C is the cumulated oxygen path of the cloud layer ($Z_C = Z_B - Z_T$).

155 mations from RSS measurements. We will apply the same
156 algorithm for one-year data at the ARM SGP site.

157 2.2. Detection Method

158 [7] In a single-layer dense cloud with fixed physical
159 depth, the photon path length scales linearly with optical
160 depth, illustrating characteristics of classic Brownian diffu-
161 sion with Gaussian statistics [Min *et al.*, 2001]. For a mul-
162 tilayered or complex cloud, a simple linear scaling does not
163 exist. In the frame of photon diffusion theory, Davis and
164 Marshak [2002] derived a mean-variance relation for a
165 homogeneous media. As shown in the study by Min *et al.*
166 [2004], the mean-variance curve with respect to a homo-
167 geneous model prediction provides a lower envelope on the
168 observed data. It demonstrated the bias of 1-D theoretical
169 calculation with respect to the more complicated 3-D
170 observation. Such characteristics, therefore, provide a diag-
171 nostic tool of 3-D scattering and absorption structures in
172 complex cloud systems. Our objective is to detect possible
173 “missed” clouds, i.e., to flag possible multilayer or complex
174 clouds that are detected by MMCR-MPL as single-layer
175 clouds. Therefore, our detection strategy is (1) to estimate
176 photon path information from the observed single-layer
177 cloud structure of MMCR-MPL and optical properties
178 retrieved from the Multifilter Rotating Shadowband Radi-
179 ometer (MFRSR), based on 1-D diffusion theory; and (2) to
180 detect the “missed” clouds by evaluating the estimated
181 photon path information against observed photon path length
182 information from a co-located RSS.

183 [8] For a single-layer cloud, sketched in Figure 2, the
184 photon path length can be separated into three intrinsically
185 linked parts: (1) transmitting from the top of the atmosphere
186 to the cloud top, (2) scattering through the cloud layer, and
187 (3) bouncing between the cloud base and the surface. The
188 cloud geometry, i.e., the cloud top height (H_T) and the cloud
189 base height (H_B) are determined by MMCR-MPL, whereas
190 cloud optical depth is inferred from measurements from the
191 MFRSR [Min and Harrison, 1996]. Since the photon path
192 length observed through oxygen A band measurement is a
193 pressure-weighted oxygen cumulated path length, we defined

the atmosphere and cloud geometry in terms of pressure-
194 weighted oxygen cumulated path length, i.e., Z_A , Z_B , Z_C , and
195 Z_T in Figure 2.

[9] To derive a simple baseline model for mean path
197 length in the atmosphere, we parameterized each portion as
198 follows:

[10] 1. Since there is not much scattering occurring above
200 the cloud layer, the path length from the top of the atmo-
201 sphere to the cloud top is simply, $M_T = Z_T/\cos(\text{SZA})$, where
202 SZA is the solar zenith angle.

[11] 2. In the diffusion limit of multiple scattering, the
204 mean path length (pressure-weighted oxygen mean path
205 length) within the cloud layer is proportional to the product
206 of cloud thickness Z_C (pressure-weighted oxygen cumulated
207 path length in cloud) and vertical cloud optical depth $Z_C\tau$,
208 since the total number of scatterings for transmitted photons
209 N is proportional to τ^2 and the total path length $M = \text{mfp}N =$
210 $(H/\tau)\tau^2 = H\tau$ [Davis and Marshak, 2002]. Because of the
211 photon penetration for the first scattering, the first scattering
212 path length is sensitive to the location of the cloud top (Z_T is
213 pressure-weighted oxygen cumulated path length from the
214 top of cloud to the top of atmosphere) and solar zenith angle.
215 Therefore, the total mean path length within the cloud layer
216 can be expressed as $M_C = Z_C(c_1 + c_2\tau + c_3Z_T/(Z_A * \cos$
217 $(\text{SZA})))$.

[12] 3. The mean path length due to the bounce between
219 the cloud base and the surface can be assumed as $M_B =$
220 $(Z_A - Z_B)\tau^{c_4}$, as cloud reflection is related to cloud optical
221 depth.

[13] Therefore, the mean path length in the atmosphere for
223 a single-layer cloud can be parameterized as

$$M = M_T + M_C + M_B = Z_T/\cos(\text{SZA}) + Z_C(c_1 + c_2\tau + c_3Z_T/(Z_A * \cos(\text{SZA}))) + (Z_A - Z_B)\tau^{c_4}.$$

The variance of photon path length is proportion to the
225 square of the product of cloud geometric thickness and
226 optical depth in diffusion limit [Davis and Marshak, 2002].
227 Similar to the mean path length, a simple model for variance
228 is also developed as $\text{var} = p_1/\cos(\text{SZA})^2 + p_3Z_C^2\tau^2 + (Z_A -$
229 $Z_B)^2\tau^{p_4}$, where c_1 , c_2 , c_3 , c_4 , p_1 , p_2 , p_3 , and p_4 are coeffi-
230 cients to be determined in the real atmosphere. To evaluate
231 this parameterization and determine those coefficients, we
232 used a Monte Carlo radiative transfer model to simulate
233 thousands of cloud fields and associated photon path length
234 distributions, including single-layer and multilayer clouds
235 with various cloud locations, cloud thicknesses, and cloud
236 optical depths. For single-layer clouds, we set cloud optical
237 depth varying from 10 to 80, cloud base from 0 to 8 km,
238 cloud thickness from 0.5 to 6 km, and solar zenith angle
239 from 0° to 70° . For multilayer clouds, we added additional
240 cloud layers above previous simulated single-layer clouds
241 with different cloud properties. Although thousands of cloud
242 fields may not include all possible cloud scenarios in the real
243 atmosphere, they provide a basic set for understanding the
244 relationship between photon path length information and
245 cloud physical and optical properties, in terms of differen-
246 tiating single-layer clouds from multilayer clouds.

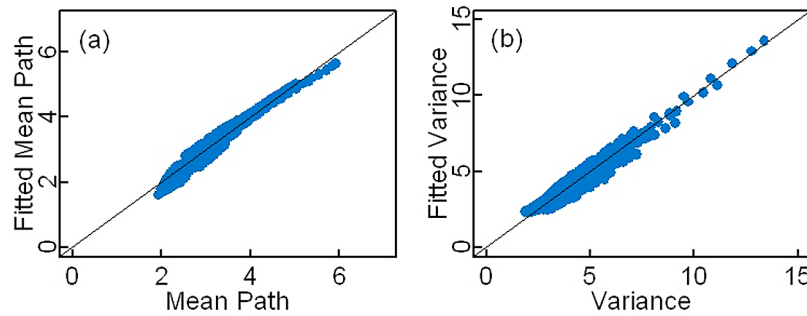


Figure 3. Fitted mean and variance compared to Monte Carlo radiative model-simulated mean and variance.

248 [14] The above simple parameterizations provide estima-
 249 tions of mean and variance of photon path length distribu-
 250 tion for single-layer clouds, using the cloud geometric
 251 and optical properties observed from MMCR-MPL and
 252 MFRSR. Figure 3 shows the comparison of simulated and
 253 fitted mean and variance of photon path length distribution
 254 based on Monte Carlo simulations of single-layer clouds. The
 255 maximum differences between the simulated and fitted mean
 256 and variance are 0.5 and 1.3, respectively. Those maximum
 257 fitting errors provide detection limits for our method. For
 258 multilayer cloud systems, multiple scattering within the lay-
 259 ers and between layers will substantially enhance the photon
 260 path length. If the observed mean path length (and/or vari-
 261 ance) is much larger than the fitted mean (and/or fitted vari-
 262 ance), i.e., greater than the maximum fitting errors, we flag it
 263 as a possible multilayer cloud. Specifically, as shown in
 264 Figure 4, all the single-layer clouds are located in the corner
 265 of the joint statistics of the Δ -mean (or the mean path length
 266 difference defined as observed (or “simulated”) mean – fitting
 267 mean) and the Δ -variance (or variance difference defined
 268 as observed (or “simulated”) variance – fitted variance),
 269 which distinctly separate them from most multilayer clouds
 270 (Figure 4b). Certainly, there are some multilayer clouds
 271 with the joint statistical characteristics overlapped with
 272 single-layer clouds. Those multilayer clouds may either
 273 have too small vertical separation between the layers or
 274 have the same first two moments as single-layer clouds
 275 with different higher moments of photon path length dis-
 276 tribution. To further distinguish those multilayer clouds
 277 from single layer clouds, it requires higher resolution of

oxygen A band measurements that are able to retrieve 278
 higher moments of photon path length distribution. Given 279
 current resolution of RSS, only the first two moments can 280
 be retrieved [Min and Clothiaux, 2003]. Therefore, there 281
 are two possible thresholds for distinguishing multilayer 282
 clouds from single-layer clouds. The dashed line represents 283
 the normal thresholds, under which all single-layer clouds 284
 are included. It is determined by the maximum differences 285
 between the simulated and fitted mean and variance. Within 286
 this threshold, however, some multilayer clouds are treated 287
 as single-layer clouds. The black solid lines represent the 288
 conservative threshold, the values of which are 20% larger 289
 than the normal threshold on Δ -mean and 50% larger than 290
 the normal threshold on Δ -variance. The additional 20% 291
 and 50% in mean and variance are much more than the 292
 maximum fitting errors. Although this conservative thresh- 293
 old results in more multilayer clouds being identified as 294
 single-layer clouds, it provides the most conservative 295
 detection of possible “missed” clouds from MMCR-MPL 296
 single-layer clouds. As the diffusion theory holds for opti- 297
 cally thick clouds, only clouds with optical depth greater 298
 than 10 will be considered in the observation. 299

3. Results

300

[15] We processed the measurements of MMCR, RSS, and 301
 MFRSR at the ARM SGP site for the year 2000. The cloud 302
 boundary and layer information were based on ARSCL that 303
 combined the measurements of MMCR and MPL [Clothiaux 304
et al., 2000]. The first two moments of photon path length 305
 distribution were retrieved from the RSS, whereas the cloud 306

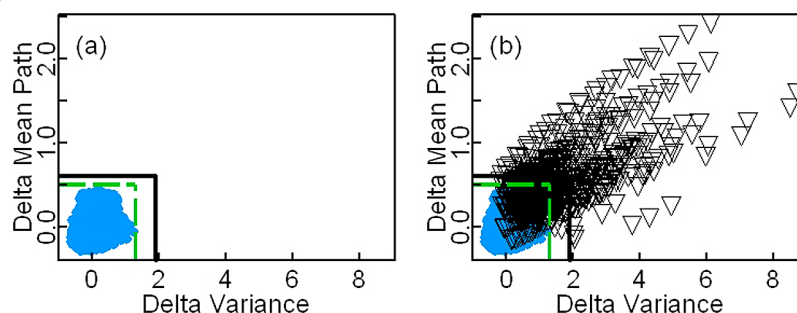


Figure 4. Δ -Mean and Δ -variance for single-layer cloud and multilayer cloud: (a) single-layer clouds (blue dots); (b) single-layer clouds (blue dots) and multilayer clouds (black triangles). The dashed green lines and solid black lines are for the normal threshold and the conservative threshold, respectively.

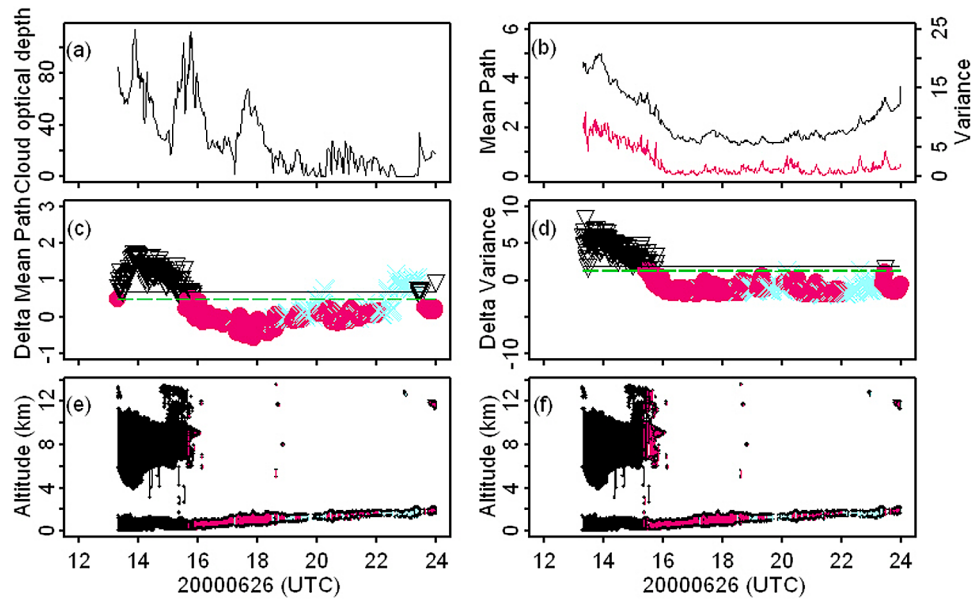


Figure 5. Time series plots. (a) Cloud optical depth retrieved from MFRSR; (b) mean path length (black line) and variance (red line) retrieved from RSS; (c) Δ -mean: the green dashed lines and black solid lines are for the normal and conservative thresholds, respectively; black triangles stand for those points over the normal threshold; (d) Δ -variance; (e) cloud profiles retrieved from MMCR-MPL with the combined normal threshold classification: black, red, and light blue colors stand for multilayer clouds, single-layer clouds, and optically thin clouds ($\tau > 10$), respectively; and (f) cloud profile classification with the combined conservative threshold.

307 optical depth was obtained from the MFRSR. Before pre-
 308 senting year-long statistics, we showed four cases to illustrate
 309 the feasibility of our detection method.

310 3.1. Case 1 (26 June 2000)

311 [16] As shown in Figure 5e, on 26 June 2000, the
 312 MMCR-MPL detected a low-level cloud persistently
 313 through the day with multilayer clouds in the morning and
 314 late in the afternoon. Retrieved cloud optical depths from the
 315 MFRSR, shown in Figure 5a, varied from very thick (over
 316 105) in the morning to very thin (less than 5) in the after-
 317 noon. Both inferred mean path length and variance from the
 318 RSS varied in concert with cloud optical depths (Figure 5b),
 319 which is consistent with our previous findings [Min et al.,
 320 2001; Min and Clothiaux, 2003]. Substantial changes in
 321 solar zenith angle or air mass cause the both mean and
 322 variance of photon path length distribution to vary in a large
 323 range. Enhancements in both the mean and variance of
 324 photon path length distribution due to multilayer clouds are
 325 relatively smaller than the changes associated with variation
 326 of solar zenith angle. Therefore, the detection power of
 327 multilayer clouds directly from the mean and variance of
 328 photon path length distribution is limited.

329 [17] After properly removing the path length contribution
 330 from the lower layer clouds as outlined in section 2, the Δ -
 331 mean and Δ -variance, shown in Figures 5c and 5d, exhibit
 332 strong distinguishing power. On the basis of the normal
 333 (dashed line) or conservative (solid line) detection thresh-
 334 olds, cloud fields can be divided into multilayer clouds
 335 (black) and single-layer clouds (red), shown in Figures 5e
 336 and 5f, respectively. Because of the limit of the diffusion

theory, optically thin clouds (optical depth < 10) are 337
 excluded from analysis and marked as light blue. Clearly, 338
 most multilayer clouds observed by MMCR-MPL were 339
 identified by the photon path length method. Some multi- 340
 layer clouds with a very thin upper layer were classified as 341
 being single layered by both thresholds. With the conser- 342
 vative threshold, more multilayer clouds were classified as 343
 single-layer clouds, as expected. This case illustrates the 344
 detection power of the photon path length method. 345

346 3.2. Case 2 (2 June 2000)

[18] The case of 2 June 2000, shown in Figure 6, was a 347
 special case where occasionally upper-level clouds appeared 348
 above a physically thick lower-level cloud deck. Because of 349
 the large thickness of the lower-level cloud, most of the 350
 photon path length was accumulated within the lower-level 351
 cloud layer. With the normal detection threshold, the Δ - 352
 mean diagnosed that this cloud system was a single-layer 353
 cloud. Even with the conservative detection threshold, the 354
 Δ -mean indicated most clouds were single-layer clouds, 355
 except for some multilayer clouds around 19:00 UTC. It 356
 suggests that enhanced path length due to the upper layer 357
 cloud was relatively small and Δ -mean is not sensitive 358
 enough for this thick low level cloud situation. However, as 359
 shown in Figure 6d, the multilayer clouds diagnosed by Δ - 360
 variance were consistent with MMCR-MPL observation 361
 (Figures 6e–6f). The difference between the normal and 362
 conservative thresholds was small. It is clear that for thick 363
 low level cloud situation, Δ -variance is more sensitive to 364
 multilayer clouds than Δ -mean. 365

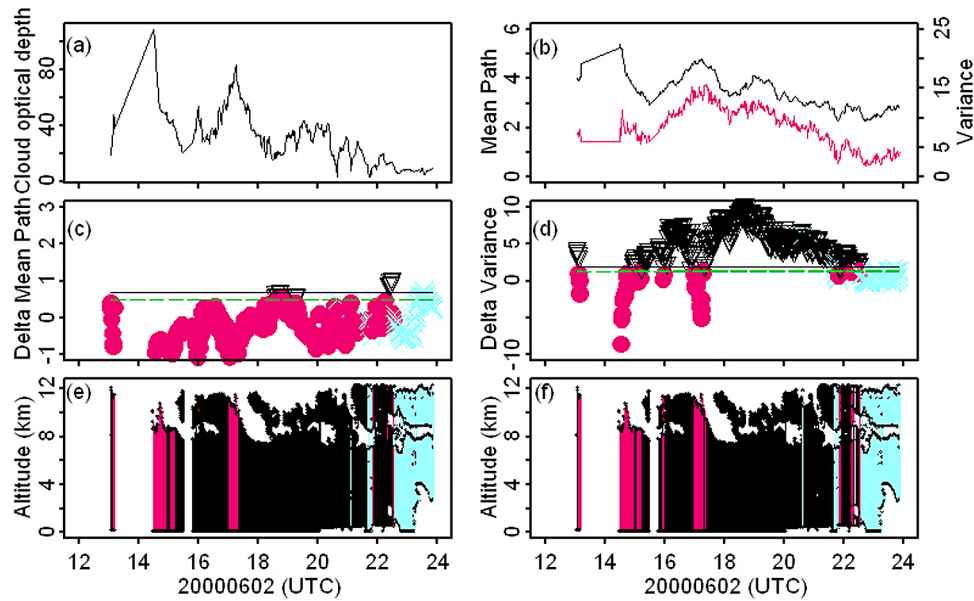


Figure 6. The same as Figure 5, but for 2 June 2000.

366 3.3. Case 3 (21 March 2000)

367 [19] Our photon path detection method based on diffusion
 368 theory is particularly good for optical thick situations. Clouds
 369 that occurred on 21 March 2000, as shown in Figure 7, were
 370 optically thick ($\tau > 30$). However, as the upper-level clouds
 371 were relatively thin compared to the lower-level clouds, our
 372 path length methods (Δ -mean and Δ -variance) classified
 373 some MMCR-MPL detected multilayer clouds as the single-
 374 layer cloud. It suggests that our detection of single-layer
 375 clouds is quite relaxed, allowing some interference of upper-
 376 level clouds. Keeping the relaxation in mind, it is interesting
 377 to see the period from 14.8 UTC and 15.7 UTC. During this

period, the MMCR-MPL detected just a single low-level 378
 cloud. However, both Δ -mean and Δ -variance with the 379
 conservative thresholds diagnosed this period as a multilayer 380
 cloud period. It means that under optically thick conditions, 381
 the radiation field, as indicated by photon path length dis- 382
 tribution, violated the diffusion theory of a single-layer 383
 cloud. In other words, the radiation field was influenced by 384
 some clouds other than the MMCR-MPL-detected clouds. 385
 Those clouds were either out of the field of view of the 386
 MMCR-MPL but within the scale of cloud-radiation inter- 387
 action or above the MMCR-MPL but having hydrometeors 388
 that were too small to be detected by the MMCR-MPL. 389

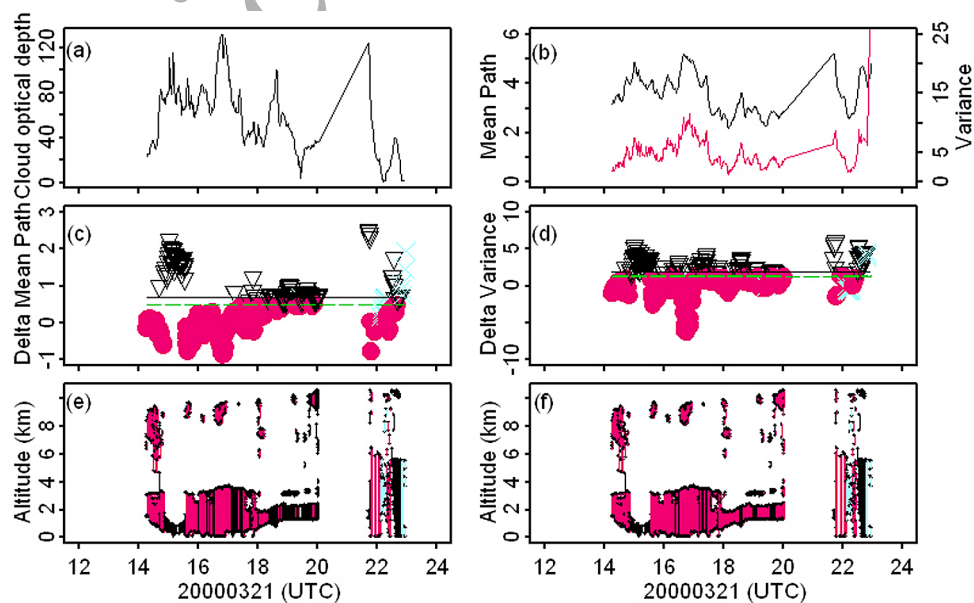


Figure 7. The same as Figure 5, but for 21 March 2000.

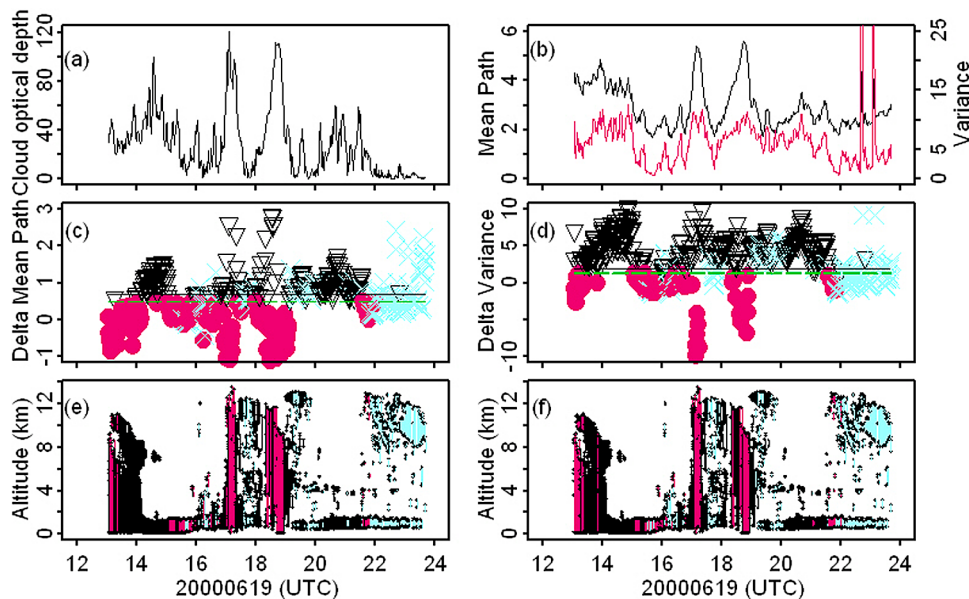


Figure 8. The same as Figure 5, but for 19 June 2000.

390 3.4. Case 4 (19 June 2000)

391 [20] The case of 19 June 2000, shown in Figure 8, is
 392 another interesting case. The clouds between 13.1 UTC and
 393 14.3 UTC were deep convective clouds with a broken layer
 394 in the early morning. Those deep convective clouds
 395 occurred again at 17.5 UTC and late around 18.7 UTC. For
 396 the rest of the time, a low-level cloud persisted with occa-
 397 sionally scattered upper-level clouds. As shown in Figures
 398 8c and 8d, under physically thick cloud conditions, the Δ -
 399 variance is more sensitive to diagnose multilayer clouds
 400 than the Δ -mean, which further corroborates the finding in
 401 case 2. The cloud field classification from the photon path
 402 length method is very consistent with the MMCR-MPL
 403 observation except for a few periods.

404 [21] Within the period of 14.3–15.1 UTC, both Δ -mean
 405 and Δ -variance diagnosed the clouds as being multilayered,
 406 whereas the MMCR-MPL detected only two scattered
 407 upper-level clouds around 14.6 UTC and 14.7 UTC. It could
 408 be either the 3-D effect of scattered upper-level clouds
 409 impacted the nearby radiation field or some other clouds
 410 existed but were not detected by the MMCR-MPL. A
 411 similar situation occurred for the period of 20.0–21.7 UTC.
 412 More interestingly, for the period of 15.1–15.6 UTC, both
 413 photon path length method and the MMCR-MPL detected a
 414 single-layer cloud, except for the period between 15.4 and
 415 15.5 UTC. During this 6 min interval, both Δ -mean and

Δ -variance diagnosed the clouds as multilayer clouds. It
 416 could be the situation that a cloud was aloft somewhere
 417 but beyond the field of view (FOV) of the MMCR-MPL.
 418

4. Aggregate Statistics and Sensitivity Study 419

[22] The case studies provide some insights on how the
 420 photon path length method works for diagnosing multilayer
 421 clouds. It is important to assess possible “missed” clouds
 422 by the MMCR-MPL statistically. We applied this method
 423 to 1 year (year 2000) measurements at the ARM SGP site.
 424 Over 59% of all clouds (daytime and nighttime) were
 425 detected by MMCR-MPL as single-layer clouds, whereas
 426 about 34% of all clouds occurred in the daytime with solar
 427 zenith angles less than 70° . Most clouds during the day-
 428 time were optically thin clouds, and only 32.2% of those
 429 single-layer clouds were optically thick ($\tau > 10$). About
 430 56% of those optically thick clouds were detected by the
 431 MMCR-MPL as single-layer clouds.
 432

[23] As listed in Table 1, with the normal threshold, the
 433 consistency rate between the photon path length method and
 434 the MMCR-MPL detection were 66.5% and 56.4% for
 435 single-layer clouds and multilayer clouds, respectively. It
 436 means that with the normal threshold the photon path length
 437 method diagnosed 43.6% of multilayer clouds as being
 438 single layered. In the meantime, about 33.5% of the
 439 MMCR-MPL detected single-layer clouds were diagnosed
 440

t1.1 **Table 1.** Aggregate Statistic Under the Normal Threshold for the Year 2000 at the ARM SGP Site^a

t1.2	Normal Threshold	MMCR-MPL Single-Layer Cloud	MMCR-MPL Multilayer Cloud
t1.3	A band single-layer cloud	66.5% (35.8%)	43.6% (20.1%)
t1.4	A band multilayer cloud	33.5% (18.0%)	56.4% (26.1%)

t1.5 ^aThe values outside parentheses are the percentages of A band detection over analyzed MMCR-MPL detection (with solar
 t1.6 zenith angle less than 70° and optical depth larger than 10), whereas the values in parentheses are the percentages of A band
 t1.7 detection over all analyzed clouds.

t2.1 **Table 2.** Same as Table 1 But Under the Conservative Threshold

t2.2	Conservative Threshold	MMCR-MPL Single-Layer Cloud	MMCR-MPL Multilayer Cloud
t2.3	A band single-layer cloud	72.3% (39.0%)	52.7% (24.3%)
t2.4	A band multilayer cloud	27.7% (14.9%)	47.3% (21.8%)

441 by the photon path length method as multilayer clouds. It
 442 suggests that one third of the MMCR-MPL-detected opti-
 443 cally thick single-layer clouds had been influenced radi-
 444 tively by other “missed” clouds.

445 [24] Even with the conservative threshold (Table 2) that
 446 allowed over half of the MMCR-MPL detected multilayer
 447 clouds to be classified as single-layer clouds, there were still
 448 27.7% of the MMCR-MPL detected single-layer clouds that
 449 were diagnosed by the photon path length method as mul-
 450 tilayer clouds. With this conservative estimation, at least,
 451 one quarter of the MMCR-MPL-detected single-layer
 452 clouds had been influenced by other clouds; either the
 453 clouds were composed of small hydrometeors and/or thinner
 454 than the radar sample volume depth resulting in partial beam
 455 filling or somewhere beyond the FOV of the MMCR-MPL.

456 5. Conclusion

457 [25] From the perspective of the GCM, the most important
 458 reason to do radiative calculations in any form is to obtain
 459 the broadband heating rates. As the BBHRP products in the
 460 ARM program primarily use cloud products from the
 461 MMCR-MPL, “missed” cloud layers in current MMCR-
 462 MPL retrievals result in substantial errors in the BBHRP
 463 products. To flag those potential multilayer clouds “missed”
 464 by MMCR-MPL, we developed a detection method based
 465 on photon path length distribution. Our photon path length
 466 method is to estimate photon path length information from
 467 the low-level single-layer cloud structure that can be accu-
 468 rately observed by the MMCR-MPL and optical properties
 469 from the MFRSR and to detect the “missed” clouds. As
 470 multiple scattering within the cloud layers and between
 471 layers would substantially enhance the photon path length,
 472 the multilayer clouds can be diagnosed by evaluating the
 473 estimated photon path information against observed photon
 474 path length information from a co-located RSS. Using a
 475 Monte Carlo radiative transfer model, we parameterized both
 476 mean and variance of the photon path length distribution for
 477 single-layer cloud structure, based on the classic diffusion
 478 theory. The maximum errors between the simulated and
 479 fitted mean and variance were 0.5 and 1.3, respectively.
 480 Those maximum fitting errors provide a measure of detec-
 481 tion uncertainty in both Δ -mean and Δ -variance schemes.

482 [26] We processed the measurements of MMCR-MPL,
 483 RSS, and MFRSR at the ARM SGP site for the year 2000.
 484 Cases studies illustrated the consistency between MMCR-
 485 MPL detection and the photon path length method under
 486 most conditions. Also for the thick, low-level clouds, Δ -
 487 variance is more sensitive to diagnose the multilayer clouds
 488 than Δ -mean. Even with both normal and conservative
 489 thresholds that allow some multilayer clouds to be diag-
 490 nosed as single-layer clouds, the photon path length method
 491 detected some multilayer clouds that were detected by the
 492 MMCR-MPL as single-layer clouds. It means that the upper

layer clouds “missed” by the MMCR-MPL had significant
 effects on radiation, e.g., photon path length. On the basis of
 1 year statistics at the ARM SGP site, we found that about
 27.7% of single-layer clouds detected by the MMCR-MPL
 with solar zenith angle less than 70° and optical depth
 greater than 10 could be multilayer clouds. It is a conser-
 vative estimation with the conservative threshold that treats
 over half of the MMCR-MPL detected multilayer clouds to
 be classified as single-layer clouds.

[27] Our photon path length method has some limitations.
 It is based on a passive instrument, which is only applicable
 during daytime. Also, our parameterization of both mean
 and variance is based on diffusion theory with optically
 thick assumption. Nonetheless, within the detection limits,
 the photon path length method diagnosed over 27% of the
 MMCR-MPL detected single-layer clouds could be influ-
 enced radiatively by other “missed” clouds. We should flag
 those periods and be cautious of any radiation application of
 the MMCR-MPL measurements during those periods. Fur-
 thermore, under other conditions, optically thin clouds or
 clouds that occurred during nighttime, we suspect that a
 substantial portion of single-layer clouds detected by the
 MMCR-MPL could also be influenced by some “missed”
 clouds or by the 3-D effects of clouds. Without accurately
 detecting those “missed” clouds, the BBHRP will be inac-
 curate. Our results echo the need for a true 3-D scanning
 radar for radiation applications. Also, our photon path length
 information is retrieved from the modest resolution mea-
 surements of RSS. Only the first two moments (mean and
 variance) of photon path length distribution can be inferred,
 which further limits our detection capability of 3-D cloud
 effects. With a high-resolution oxygen A band spectrometer
 [Min et al., 2004], we expect a more powerful diagnosis for
 3-D cloud effects from retrieved higher moments of photon
 path length distribution.

References

- Clothiaux, E. E., T. P. Ackerman, G. G. Mace, K. P. Moran, R. T. Marchand,
 M. A. Miller, B. E. Martner (2000), Objective determination of cloud
 heights and radar reflectivities using a combination of active remote sen-
 sors at the ARM CART sites, *J. Appl. Meteorol.*, 39, 645.
 Davis, A. B., and A. Marshak (2002), Space-time characteristics of light
 transmitted through dense clouds: A Green’s function analysis, *J. Atmos.*
Sci., 59, 2713–2727.
 Fischer, J., and H. Grassl (1991), Detection of cloud-top height from back-
 scattered radiances within the oxygen A band, Part 1: Theoretical study,
J. Appl. Meteorol., 30, 1245.
 Fischer, J., W. Cordes, A. Schmitz-Peiffer, W. Renger, and P. Morel
 (1991), Detection of cloud-top height from backscattered radiances
 within the oxygen A band, part 2: Measurements, *J. Appl. Meteorol.*,
 30, 1260.
 Grechko, Y. I., V. I. Dianov-Klokov, and I. P. Malkov (1973), Aircraft
 measurements of photon paths in reflection and transmission of light
 by clouds in the $0.76 \mu\text{m}$ oxygen band, *Atmos. Ocean Phys.*, 9, 262.
 Harrison, L., and Q.-L. Min (1997), Photon path length distributions in
 cloudy atmospheres from ground-based high-resolution O_2 A band spec-
 troscopy, in *IRS’96: Current Problems in Atmospheric Radiation*, edited
 by W. L. Smith and K. Stamnes, p. 594, A. Deepak, Hampton, Va.

- 550 Irvine, W. M. (1964), The formation of absorption bands and the distribu- 575
 551 tion of photon optical paths in a scattering atmosphere, *Bull. Astron. Inst.* 576
 552 *Neth.*, 17, 266–279. 577
- 553 Irvine, W. M. (1966), The shadowing effect in diffuse, *J. Geophys. Res.*, 578
 554 71(12), 2931–2937, doi:10.1029/JZ071i012p02931. 579
- 555 Min, Q.-L., and L. C. Harrison (1996), Cloud properties derived from sur- 580
 556 face MFRSR measurements and comparison with GOES results at the 581
 557 ARM SGP site, *Geophys. Res. Lett.*, 23(13), 1641, doi:10.1029/ 582
 558 96GL01488. 583
- 559 Min, Q.-L., and L. C. Harrison (1999), Joint statistics of photon path length 584
 560 and cloud optical depth, *Geophys. Res. Lett.*, 26(10), 1425, doi:10.1029/ 585
 561 1999GL900246. 586
- 562 Min, Q.-L., L. C. Harrison, and E. E. Clothiaux (2001), Joint statistics of 587
 563 photon path length and cloud optical depth: Case studies, *J. Geophys.* 588
 564 *Res.*, 106(D7), 7375, doi:10.1029/2000JD900490. 589
- 565 Min, Q.-L., and E. E. Clothiaux (2003), Photon path length distributions 590
 566 inferred from rotating shadowband spectrometer measurements at the 591
 567 atmospheric Radiation Measurements Program Southern Great Plains 592
 568 site, *J. Geophys. Res.*, 108(D15), 4465, doi:10.1029/2002JD002963. 593
- 569 Min, Q.-L., and L. C. Harrison (2004), Retrieval of atmospheric optical 594
 570 depth profiles from downward-looking high-resolution O₂ A band mea- 595
 571 surements: Optically thin conditions, *J. Atmos. Sci.*, 61, 2469–2477. 596
- 572 Min, Q.-L., L. C. Harrison, P. Kiedron, J. Berndt, and E. Joseph (2004),
 573 A high-resolution oxygen A band and water vapor band spectrometer,
 574 *J. Geophys. Res.*, 109, D02202, doi:10.1029/2003JD003540.
- O’Brian, D. M., and R. M. Mitchell (1992), Error estimates for retrieval of 575
 cloud top pressure using absorption in the A band of oxygen, *J. Appl.* 576
Meteorol., 31, 1179. 577
- Pfeilsticker, K., F. Erle, H. Veitel, and U. Platt (1998), First geometrical 578
 path lengths probability density function derivation of the skylight 579
 from spectroscopically highly resolving oxygen A band observations, 580
 1: Measurement technique, atmospheric observations and model calcu- 581
 lations, *J. Geophys. Res.*, 103(D10), 11,483–11,504, doi:10.1029/ 582
 98JD00725. 583
- Portmann, R. W., S. Solomon, R. W. Sanders, and J. S. Danel (2001), 584
 Cloud modulation of zenith sky oxygen path lengths over Voulder, 585
 Colorado: Measurements versus model, *J. Geophys. Res.*, 106(D1), 586
 1139, doi:10.1029/2000JD900523. 587
- Van de Hulst, H. C. (1980), *Multiple Light Scattering, Tables, Formulas* 588
and Applications, vol. 1 and 2, Academic, London. 589
- Veitel, H., O. Funk, C. Kruz, U. Platt, and K. Pfeilsticker (1998), Geomet- 590
 rical path length probability density functions of the skylight transmitted 591
 by midlatitude cloudy skies: Some case studies, *Geophys. Res. Lett.*, 25 592
 (17), 3355, doi:10.1029/98GL02506. 593
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